



# Eddy Covariance Fluxes of NO and NO<sub>2</sub> above a Northeastern U.S. Forest

## I. Abstract

Although tropospheric NO, NO<sub>2</sub>, and O<sub>3</sub> rapidly interconvert in a fast photochemical cycle, the processes governing their removal rates, interactions with ecosystems, and human interest in their mitigation are distinctly different. Thus a detailed understanding of their behavior at the atmosphere-biosphere interface is crucial. The reactivity of this family of gases has often confounded measurements in the past, and the question of biospheric emission and uptake of NO<sub>x</sub> remains open (Lerdau et al., 2000).

In order to address these issues, we present concurrent eddy covariance fluxes of NO, NO<sub>2</sub>, and O<sub>3</sub> at the rural, deciduous, mixed hardwood Harvard Forest in central Massachusetts during the summer and fall of 2000. The independent measurements were conducted above the forest canopy, at similar heights, on two towers within several hundred meters of one another. NO was measured using an existing photolysis-chemiluminescence detector in the eddy covariance mode, and NO<sub>2</sub> using a new tunable diode laser absorption spectrometer (TDLAS). O<sub>3</sub> concentrations, profiles, and fluxes were also measured at the site. At night, NO fluxes were effectively zero while small but persistent downward NO<sub>2</sub> fluxes were observed. Fluxes of NO<sub>2</sub> during the day were generally upward and coupled with downward NO fluxes of similar magnitude. The opposing NO and NO<sub>2</sub> fluxes confirm the predictions of some canopy exchange models that include the effect of the light gradient within the canopy on photo-reactive species (e.g. Gau et al. 1993).



Figure 2. Left: Inlet and sonic anemometer at a height of 22m. The canopy height averages 20 m. Right: View of Harvard Forest with sampling tower location marked.

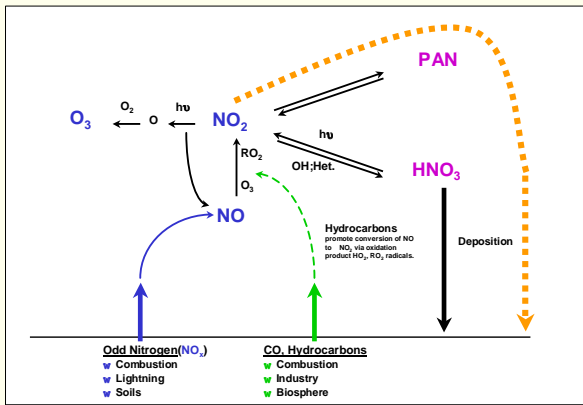


Figure 1. Simplified schematic of tropospheric NO-NO<sub>2</sub>-O<sub>3</sub> and reactive nitrogen chemistry.

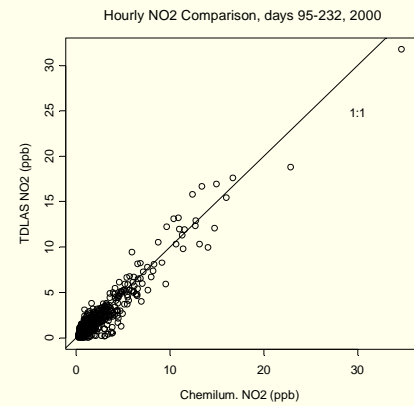


Figure 3. Hourly TDLAS and photolysis-chemiluminescence NO<sub>2</sub> concentrations at Harvard Forest. The two instruments operated from separate towers roughly 200 m apart at 22m (TDLAS) and 29m (chemilum.) sampling heights. Orthogonal distance fit  $y=a+bx$  (errors at 95<sup>th</sup> confidence interval):  $a = -0.2 \pm 1.0$ ,  $b = 1.1 \pm 0.2$ ,  $R^2 = 0.91$ .

## II. Methods

### INSTRUMENTATION:

From late August to mid-October, an existing chemiluminescence detector (29 m sampling height) was configured to measure NO concentrations at 8 Hz. At other times, the chemiluminescence detector measured slow profiles and concentrations of NO and NO<sub>2</sub>, converted to NO in a photolysis cell prior to detection. We installed a new tunable diode laser absorption spectrometer (TDLAS) on a nearby tower to measure NO<sub>2</sub> (22 m sampling height) at 8 Hz. The TDLAS concentration measurements were verifiably species-specific, spectroscopically calibrated, and compared very well to ongoing photolysis-chemiluminescence NO<sub>2</sub> measurements at the site (figure 3). For further TDLAS details, refer to accompanying poster A51F-0132 and references (Horie et al. 1999). Continuing O<sub>3</sub> concentration and flux measurements employ C<sub>2</sub>H<sub>2</sub>-chemiluminescence (29 m sampling height) and UV absorbance instruments, respectively. Three-axis sonic anemometers facing into the prevailing wind direction (west) at the sampling heights on both towers provided the vertical and horizontal wind velocities and virtual temperatures needed to compute eddy covariance fluxes of heat, momentum, NO, NO<sub>2</sub>, and O<sub>3</sub> (Munger et al. 1996, 1998).

### EDDY COVARIANCE FLUXES:

We compute 30-minute fluxes from the covariance of detrended vertical wind velocity ( $w'$ ) with fluctuations of detrended temperature ( $T'$ ) or detrended trace gas concentration ( $C'$ ). Compared to those of temperature and vertical wind speed ( $w$ ), lagged correlations and cospectra for NO<sub>2</sub> exhibit expected offset and smearing due to instrument lag time and response function (figure 4). Small corrections of order 20% are routinely applied to the NO<sub>2</sub> eddy covariance fluxes to account for these instrumental properties.

### HARVARD FOREST:

The Harvard Forest site in central Massachusetts (42.54N, 72.18W; elevation, 340 m) is a 50- to 70-year old mixed deciduous forest consisting primarily of red oak and red maple, with scattered hemlock, red pine, and white pine stands. Carbon exchange, meteorology, trace gas concentrations, and eddy covariance fluxes have been measured at the site for over a decade.

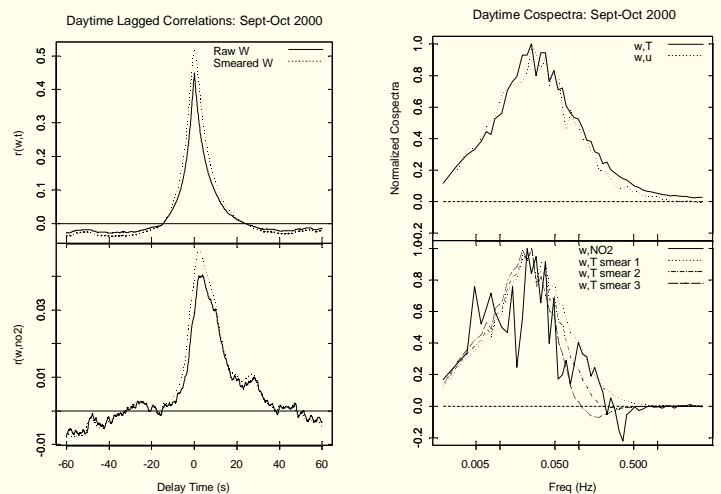
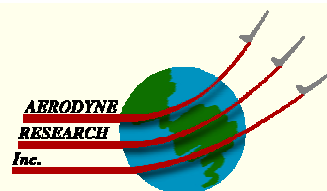


Figure 4. Average daytime lagged correlations and cospectra for ( $w,T$ ) and ( $w,NO_2$ ).  $R(w,NO_2)$  is lagged compared to  $r(w,T)$  due to transit time in the inlet, and the peak is rounded due to an exponential instrument response function with time constant 1 second. The cospectrum of  $w$  with NO<sub>2</sub> confirms that the response function is as expected when compared with cospectra of  $w$  with  $T$  smeared by 1s, 2s, and 3s exponentials.

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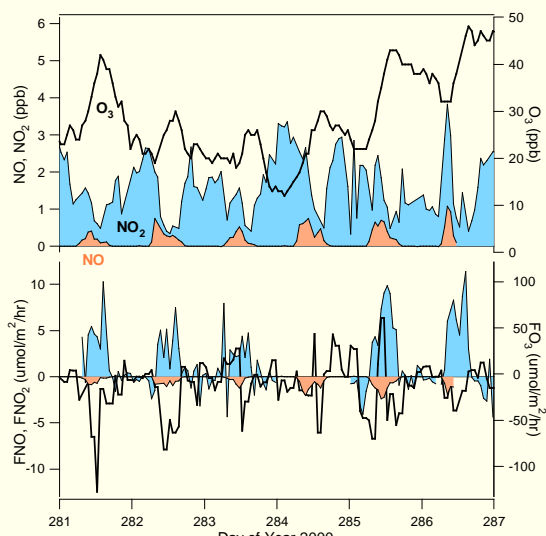
<sup>2</sup>Aerodyne Research, Inc.: Center for Atmos. and Environ. Chem., 781-663-9500, mz@aerodyne.com, http://www.aerodyne.com/caec.html



### III. Results

#### TIME SERIES

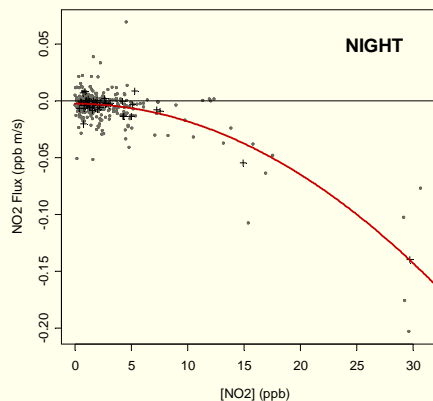
A typical Harvard Forest NO<sub>x</sub> and O<sub>3</sub> time series shows photochemical production of NO during the day and conversion to NO<sub>2</sub> at night. Daytime coupled fluxes of NO<sub>2</sub> (upward) and NO (downward) arise from photochemical cycling and turbulent transport in the presence of a forest canopy-influenced light gradient. Higher light above the canopy favors production of NO (NO<sub>2</sub>+hv → NO+O); lower light below favors conversion back to NO<sub>2</sub> (NO+O<sub>3</sub> → NO<sub>2</sub>+O<sub>2</sub>). Although the daytime fluxes appear imbalanced, the difference in measurement height of NO (29 m) and NO<sub>2</sub> (22 m) implies a flux ratio close to that observed. O<sub>3</sub> concentrations and fluxes at the site are typically an order of magnitude greater than those of NO<sub>x</sub>.



**Figure 5.** Time series of hourly NO, NO<sub>2</sub>, and O<sub>3</sub> concentrations (upper panel) and fluxes (lower panel), Oct 7-12, 2000. NO<sub>2</sub> measurements were taken at a sampling height of 22 m, NO and O<sub>3</sub> at 29 m at a nearby tower.

#### NIGHT

At night, NO concentration and flux fall to zero and the NO<sub>2</sub> flux captures total NO<sub>x</sub> deposition. The observed nighttime NO<sub>2</sub> flux was generally small and tended to be downward. An apparent quadratic dependence of NO<sub>2</sub> flux on NO<sub>2</sub> concentration could imply deposition via formation of the NO<sub>2</sub> dimer on surfaces, and is also similar to the dependence of observed nighttime HONO emissions on NO<sub>2</sub> concentration (Harrison et al., 1996). If the NO<sub>2</sub> flux represents net deposition of NO<sub>x</sub> at night, the process would depart from standard model parameterizations, especially for higher NO<sub>2</sub> concentrations.

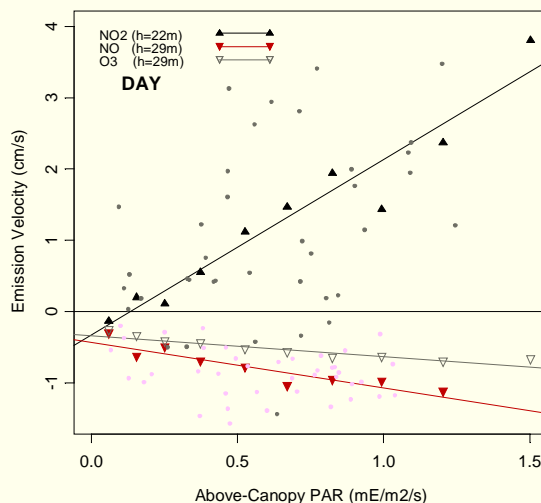


**Figure 6.** Nighttime mean hourly (dots) and nightly (pluses) NO<sub>2</sub> flux vs. concentration. A negative slope represents deposition velocity of NO<sub>2</sub>. Quadratic fit to hourly data: Flux ~ -0.00247 + -1.6e-04 \* [NO<sub>2</sub>]<sup>2</sup> (R<sup>2</sup> = 0.6). V<sub>d</sub>(NO<sub>2</sub>) appears to increase from ~0.02 cm/s at 1 ppb to ~0.4 cm/s at 25 ppb.

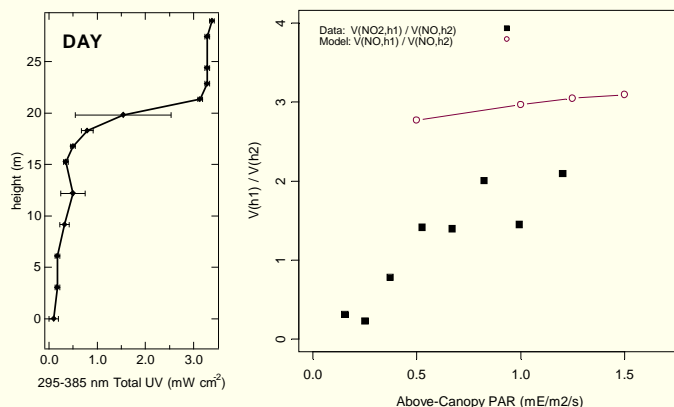
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#### DAY



**Figure 7.** Daytime emission velocities ( $V = \text{Flux}/\text{Concentration}$ ) for NO<sub>2</sub>, NO, and O<sub>3</sub> plotted as medians of hourly data within quantiles of measured above-canopy photosynthetically-active radiation (PAR). Daily median values of NO and NO<sub>2</sub> velocities are plotted as small dots. V(NO<sub>2</sub>) data span spring through fall 2000, while V(NO) data were collected late August to early October 2000 only. Linear fits to block median points:  $V(\text{NO}_2) = -0.33 + 2.5 * \text{PAR}$ , R<sup>2</sup> = 0.93.  $V(\text{NO}) = -0.43 - 0.64 * \text{PAR}$ , R<sup>2</sup> = 0.84.  $V(\text{O}_3) = -0.34 - 0.29 * \text{PAR}$ , R<sup>2</sup> = 0.85



**Figure 8.** Left: UV profile (Oct 4, 2001, local noon, clear sky); error bars at the measurement height. Right: Ratio of block median daytime NO<sub>2</sub> emission velocities (measured at h<sub>1</sub>=22m) to NO deposition velocities (measured at h<sub>2</sub>=29m) as a function of above-canopy PAR. Also shown are outputs from a simple model of NO<sub>x</sub> photochemical cycling using typical NO and NO<sub>2</sub> lifetimes, and UV light and eddy diffusivity profiles. The model solves the continuity equation for NO at each of 100 levels below and above the canopy height, assuming no net emission or deposition of NO<sub>x</sub>,  $d/dz(-K_e(d\text{NO}/dz)) = P_{\text{NO}} - L_{\text{NO}}$ , where  $K_e$  is eddy diffusivity,  $P_{\text{NO}}$  is photochemical production (NO<sub>2</sub>+hv → NO+O), and  $L_{\text{NO}}$  is chemical loss (NO+O<sub>3</sub> → NO<sub>2</sub>+O<sub>2</sub>). The over-prediction of the velocity ratio in the model is consistent with net deposition of NO<sub>x</sub> to the forest in the observed data.

### IV. Conclusions

- DAY: Coupled NO and NO<sub>2</sub> fluxes** in opposite directions are consistent with canopy light gradient-driven photochemical cycling of NO<sub>x</sub> with some net deposition.
- NIGHT: NO<sub>x</sub> deposition depends quadratically on [NO<sub>2</sub>]** and is on average non-zero. Deposition via non-stomatal pathways or to non-foliar surfaces may be more important than previously thought. Heterogeneous production of NO<sub>2</sub> dimer and/or HONO may also play a role in deposition.